Due: 18 Oct. 2018

You are welcome (in fact, encouraged) to discuss problem sets with your classmates. However, each of you must turn in your own problem-set answers.

1. In z^* coordinates, the change of geopotential Φ with height is proportional to temperature T following

$$\frac{\partial \Phi}{\partial z^*} = RT/H$$

Using this and the definition of z^* = - H ln(p/p_o), transform the pressure-coordinate thermodynamic equation

$$\frac{\partial T}{\partial t} + u \frac{\partial T}{\partial x} + v \frac{\partial T}{\partial y} - S_p \omega = \frac{J}{C_p},$$

where

$$S_p \equiv \frac{RT}{C_p} - \frac{\partial T}{\partial p} = -\frac{T}{\theta} \frac{\partial \theta}{\partial p}$$
.

into the thermodynamic equation in z* coordinates:

$$\left(\frac{\partial}{\partial t} + \vec{V} \cdot \nabla\right) \frac{\partial \Phi}{\partial z^*} + w^* N^2 = \kappa J/H,$$

where
$$N^2=rac{R}{H}\Big(rac{\partial T}{\partial z^*}+rac{\kappa T}{H}\Big)$$
 .

Here R is the gas constant, H a scale height, w* is the vertical motion in z* coordinates, J is diabatic heating, $K = R/C_p$ and C_p is the heat capacity of air at constant pressure.

2. (Holton's problem 12.1)

Suppse that temperature increases linearly with height in the layer between 20 and 50 km at a rate of 2 K/km. If the temperature is 200K at 20 km, find the value of the scale height H for which the log-pressure height z^* coincides with the actual height z at 50 km. You should assume that z^* coincides with the actual height at 20 km and that $g = 9.81 \text{ m/s}^2$ at all levels (i.e.,, it is a constant).