

Derivation of Coriolis force

To derive the zonal component, $\left(\frac{\partial u}{\partial t}\right)_\theta$, consider a latitudinal or vertical displacement and utilize conservation of angular momentum.

Equate the total angular momentum at an initial distance R to a displaced distance $R + \delta R$. (if R changes U must also change).

$$\left(\Omega + \frac{U}{R}\right) R^2 = \left(\Omega + \frac{U + \delta U}{R + \delta R}\right) (R + \delta R)^2$$

- Solve for δU : Expand RHS, neglect 2nd order differentials

$$\left(\Omega + \frac{U}{R}\right) R^2 = \left(\Omega + \frac{U + \delta U}{R + \delta R}\right) \left[R^2 + 2R\delta R + \cancel{(\delta R)^2}^0 \right]$$

$$\left(\Omega + \frac{U}{R}\right) R^2 = R^2 \Omega + 2\Omega R\delta R + R^2 \left(\frac{U + \delta U}{R + \delta R}\right) + 2R\delta R \left(\frac{U + \delta U}{R + \delta R}\right)$$

divide through by R^2 ----

$$\cancel{\Omega} + \frac{U}{R} = \cancel{\Omega} + \frac{2\Omega \delta R}{R} + \frac{U + \delta U}{R + \delta R} + \frac{2\delta R}{R} \left(\frac{U + \delta U}{R + \delta R}\right)$$

multiply through by $R + \delta R$ ----

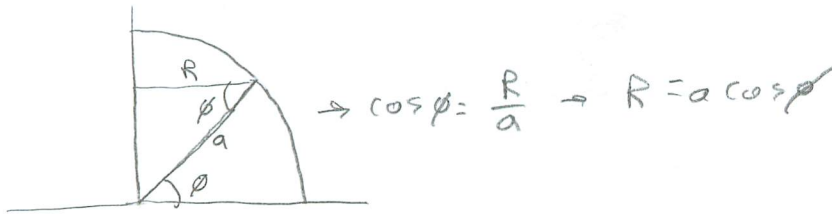
$$\cancel{U} + \frac{U\delta R}{R} = \frac{2\Omega \delta R}{R} (R + \delta R) + \cancel{U} + \delta U + \frac{2\delta R}{R} (U + \delta U)$$

$$\frac{U\delta R}{R} = 2\Omega \delta R + \cancel{\frac{2\Omega (\delta R)^2}{R}}^0 + \delta U + \frac{2U\delta R}{R} + \cancel{\frac{2\delta R \delta U}{R}}^0$$

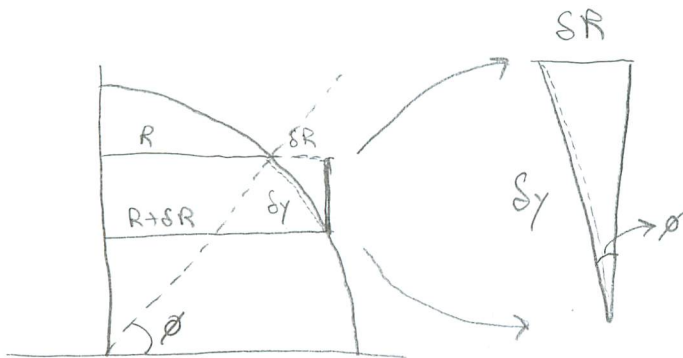
$$\frac{U \delta R}{R} = 2 \Omega \delta R + \delta U + \frac{2U \delta R}{R}$$

$$\delta U = -2 \Omega \delta R - \frac{U}{R} \delta R$$

- note that $R = a \cos \phi$



- also note that $\delta R = -\sin \phi \delta y$



it can be shown that this angle = ϕ if we take the limit as $\delta R \rightarrow 0$

$$\text{so, } \sin \phi = \frac{\delta R}{\delta y} \rightarrow \delta R = -\sin \phi \delta y$$

substitute $R = a \cos \phi$ and $\delta R = -\sin \phi \delta y$ into the equation for δU , above.

$$\delta U = +2 \Omega \sin \phi \delta y + \frac{U}{a \cos \phi} \sin \phi \delta y$$

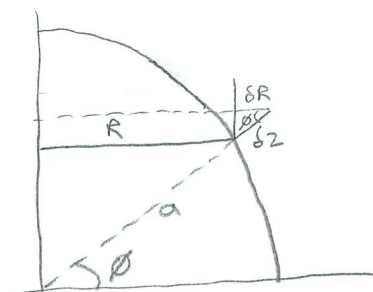
divide through by δt and take the limit as $\delta t \rightarrow 0$

$$\frac{du}{dt} = \left(2 \Omega \sin \phi + \frac{u}{a} \tan \phi \right) \frac{dy}{dt}$$

substitute $\rightarrow \frac{dy}{dt} = v$

$$\left(\frac{du}{dt} \right)_{co} = 2 \Omega v \sin \phi + \frac{uv}{a} \tan \phi$$

- For a vertical displacement $SR = \cos \phi \delta z$



$$\cos \phi = \frac{SR}{\delta z} \rightarrow SR = \cos \phi \delta z$$

- substitute $SR = \cos \phi \delta z$ into our previous equation for δu , divide by δt and take the limit as $\delta t \rightarrow 0$.

$$\frac{\delta u}{\delta t} = -2 \Omega \cos \phi \frac{\delta z}{\delta t} - \frac{u}{a \cos \phi} \cos \phi \frac{\delta z}{\delta t}$$

using $w = \frac{\delta z}{\delta t}$ we obtain ----

$$\left(\frac{du}{dt} \right)_{co} = -2 \Omega w \cos \phi - \frac{uw}{a}$$

- * To find the meridional component of the Coriolis force, $(\frac{dv}{dt})_{co}$, consider the excess in the centrifugal force: the difference in force between an East-West moving object and an object at rest.

centrifugal force $\rightarrow \frac{d\vec{v}}{dt} = \omega^2 r$.

- for a moving object on Earth $\rightarrow \frac{d\vec{v}}{dt} = (\Omega + \frac{U}{R})^2 \vec{R}$

- for a still object $\longrightarrow \frac{d\vec{v}}{dt} = \Omega^2 \vec{R}$

subtract the two ----

$$(\Omega + \frac{U}{R})^2 \vec{R} - \Omega^2 \vec{R}$$

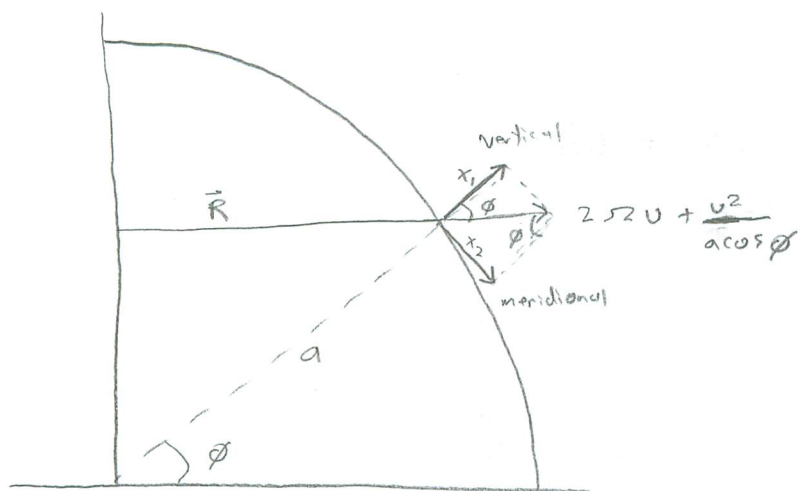
expand ----

$$\cancel{\Omega^2 \vec{R}} + \underbrace{2\Omega U \frac{\vec{R}}{R} + U^2 \frac{\vec{R}}{R^2}} - \cancel{\Omega^2 \vec{R}}$$

$\rightarrow$ these are "deflecting forces" and act in same direction as centrifugal force.

- * To find $(\frac{dv}{dt})_{co}$ for meridional and vertical displacements, take the meridional and vertical components of $2\Omega U + \frac{U^2}{R}$.

Do them one at a time ----



the vertical component,

$$\left(\frac{dw}{dt}\right)_c = 2\Omega v \cos \phi + \frac{v^2}{a}$$

the meridional component,

$$\left(\frac{dv}{dt}\right)_c = -2\Omega v \sin \phi - \frac{v^2}{a} \tan \phi$$

For synoptic scale motions $|v| \ll \Omega R$ which means that $\frac{v^2}{a} \tan \phi$ and $\frac{v^2}{a}$ become very small, thus we can approximate the Coriolis acceleration for horizontal motion as ---

$$\left(\frac{dv}{dt}\right)_c = 2\Omega v \sin \phi = f v$$

$$\left(\frac{dv}{dt}\right)_c = -2\Omega v \sin \phi = -f v$$

where $f = 2\Omega \sin \phi$ is the Coriolis parameter.