Meteorology 443

Mid-Term Test 17 April 2002

Closed-book test: 50 minutes

Point

I. The thermodynamic equation is

$$c_p \frac{dT}{dt} - \alpha \frac{dp}{dt} = \dot{q}.$$

- 5 a. Use this equation to find the relationship between θ and T. Clearly state your assumption to form this relationship.
- b. Based on the relationship between θ and T, one can obtain static stability as the following equation:

$$\frac{1}{\theta} \frac{\partial \theta}{\partial z} = \Gamma_d - \Gamma,$$

where Γ_d and Γ are dry adiabatic lapse rate and environmental lapse rate, respectively. Use the skew – T chart to explain the condition with which the atmospheric environment becomes *convectively unstable*.

II. The balanced flow without any forcing along geopotential height contours can be depicted by the equation:

$$\frac{v^2}{R} + fv = -\frac{\partial \phi}{\partial n}.$$

- What is the difference between gradient (\vec{v}_g) and geostrophic wind (\vec{v}_g) ? Use this equation to explain your answer, but you have to use the height contour structure to explain your answer.
- b. Under what synoptic condition is $|\vec{v}_g| > |\vec{v}_g|$. Explain your answer.
- 5 c. Use the gradient wind equation to explain why a regular low system has strong wind.
 - III. The thermal wind equation is

$$\frac{\partial \vec{v}_g}{\partial \ln p} = -\frac{R}{f} \, \hat{k} \times \nabla_p T$$

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- a. Use the thermal wind equation to explain why the maximum zonal wind (jet stream) exists in midlatitudes.
- 4 III. b. Why is the direction of thermal wind in the Southern Hemisphere against isothermals? Use the thermal wind equation to explain your answer.
- 10 IV. Circulation theorm vs. vorticity equation

Bjerknes circulation theorm is

$$\frac{dC}{dt} = -\oint \frac{dp}{\rho} - 2\Omega \frac{dA_e}{dt},$$

Note that $C = \xi A$ if A is small. Show that this circulation theorm is equivalent to the following Vorticity equation:

$$\frac{d\zeta}{dt} = -\nu\beta - (\zeta + f)(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y}) + \frac{1}{\rho^2}(\frac{\partial \rho}{\partial x}\frac{\partial p}{\partial y} - \frac{\partial \rho}{\partial y}\frac{\partial p}{\partial x}).$$

8 V. The surface pressure tendency equation

$$\frac{\partial p_s}{\partial t} \simeq -\int_o^{p_s} (\nabla \cdot \vec{v}) dp$$

is often used to explain the following observation:

"The development of a negative surface pressure tendency is a classic warning of an approaching cyclonic weather disturbance." The simplified vorticity equation for synoptic motions is

$$\frac{d}{dt}(\zeta + f) = -(\zeta + f) \left(\frac{\partial u}{\partial x} + \frac{\partial \mathbf{W}}{\partial y} \right).$$

Use this simplified vorticity equation to explain the aforementioned observation of the weather development caused by a cyclone disturbance.

2